## TOWARDS SNOW WATER EQUIVALENT RETRIEVAL USING INTERFEROMETRIC AND POLARIMETRIC SAR DATA

Kristina Belinska<sup>1,2</sup>, Georg Fischer<sup>1</sup>, Irena Hajnsek<sup>1,2</sup> <sup>1</sup>Microwaves and Radar Institute, German Aerospace Center <sup>2</sup>Institute of Environmental Engineering, ETH Zurich



## **Motivation**



#### **Snow Water Equivalent**

Amount of liquid water contained within a snow pack
 → depth of water, if whole snow pack melted instantaneously

$$SWE = \frac{1}{\rho_w} \int_0^{Z_s} \rho_s(z) dz \approx Z_s \rho_s / \rho_w$$

Important Parameter for

Hydrological and climate models

Water resource planning

#### Flood predictions



https://www.sieker.de/en/fachinformationen/flood/hydrologic al-modelling.html



https://www.drax.com/about-us/our-sites-andbusinesses/cruachan-power-station/



https://www.wkbw.com/news/local-news/rain-snow-melt-floods-basements-of-orchard-park-homeowners

## SAR data and test site

### SAR

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- TanDEM-X (TDX), X band
- Dual polarized data: HH and VV polarization
- Temporal baseline 11 days



#### In-situ measurements

- SWE measurements with a passive gamma ray sensor since 2019
- Temporal resolution: 6 hours



Jentzsch, K., Bornemann, N., Cable, W., Gallet, J. C., Lange, S., Westermann, S. and Boike, J. (2020): Near real-time observations of snow water equivalent for SIOS on Svalbard – (SWESOS), doi: 10.5281/zenodo.4146835



## **DInSAR model for SWE Estimation**

Repeat pass SAR acquisitions

- Different dielectric properties in snow compared to air
  - $\rightarrow$  Refraction of radar waves in the snow pack
  - → Different optical path length for snow compared to no snow conditions

Only dry snow

Snow	
Ground	



### DInSAR model for SWE Estimation



• Path delay  $\Delta R$  can be translated into an X-Band 60 C-Band interferometric phase difference L-Band 40 SWE change [mm]  $\Delta \Phi_s = 2 \; \frac{2\pi}{\lambda} \; \Delta R$ 20 0  $\Delta \Phi_s = -2 k_i \Delta Z_s (\cos \Theta - \sqrt{\epsilon - \sin^2 \Theta})$ -20 -40 $\Delta R_{air,sc}$ -60 $\epsilon_{air}$ Air -2 -3 -12  $\Delta R_{air.s}$ 0 Interferometric Phase [rad]  $\Delta R_{s,sc}$  $\epsilon_s$  $Z_s$ Snow

S. Leinss et al.: Snow Water Equivalent of Dry Snow Measured by Differential Interferometry , 2015

Ground

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## **DInSAR model for SWE Estimation**



#### Limitations

- Temporal decorrelation
- Phase calibration
- Different phase delay for different polarizations
- $\Delta \Phi_s$  between  $[-\pi, \pi] \rightarrow$  outside this interval phase wrapping errors



## SWE Estimation using DInSAR Phase

- Only limited range of SWE change can be retrieved using the X-band measurements → [-8 mm, +8 mm]
- Underestimation of SWE changes above this threshold
- In-Situ measurements used to check if SWE change lies above phase wrap threshold
  - $\rightarrow$  if yes, phase cycle is added
- Phase wraps are one of the main limitations







## **PolSAR CPD model for Snow Depth Estimation**

Additional information about snow accumulation contained

in co-polar-phase difference

$$\Phi_c = \Phi_{VV} - \Phi_{HH}$$

 Different polarizations show different propagation speeds in anisotropic snow

- Fresh snow → more horizontally aligned ice grains
  → increasing CPD values
- Decreasing CPD values due to recrystallisation under temperature gradients



 $\rightarrow$  CPD can detect snow fall events

 $\rightarrow$  help to correct phase wraps



## **PolSAR CPD model for Snow Depth Estimation**



- Snow model: ellipsoidal ice inclusions in air
- Assumption of snow anisotropy and density

 $\rightarrow$  refractive indices for HH and VV

$$\Phi_{CPD} = (-1) \frac{4\pi}{\lambda} \Sigma Z_s \left( \sqrt{n_V^2 - \sin^2(\Theta)} - \sqrt{n_H^2 - \sin^2(\Theta)} \right)$$
$$\Delta \zeta(\rho, A, \Theta)$$

Leinss et al., Snow Height Determination by Polarimetric Phase Differences in X-Band SAR Data, 2014 Leinss et al. Anisotropy of seasonal snow measured by polarimetric phase differences

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Leinss et al., Anisotropy of seasonal snow measured by polarimetric phase differences in radar time series, 2016



## **PolSAR CPD model for Snow Depth Estimation**



#### Advantages

- Less sensitive to phase wraps
- No absolute phase calibration necessary

#### Limitation for InSAR and PolSAR



**Combination of Interferometric and Polarimetric Measurements** – **Temporal Coherence region** 



- Two PolSAR acquisitions
  - $\rightarrow$  coherency matrices  $T_{11}$  and  $T_{22}$
  - $\rightarrow$  temporal PollnSAR matrix  $\boldsymbol{\Omega}_{12}$

Temporal polarimetric coherence ρ

 $\rho(\omega_1, \omega_2) = \frac{\omega_1^H \mathbf{\Omega}_{12} \omega_2}{\sqrt{(\omega_1^H \mathbf{T}_{11} \omega_1)(\omega_2^H \mathbf{T}_{22} \omega_2)}}$ 

 $\omega \rightarrow$  unitary vectors of polarization states

## TDX



Jun Ni et al., Multitemporal SAR and Polarimetric SAR Optimization and Classification: Reinterpreting Temporal Coherence, 2022

## **Combination of Interferometric and Polarimetric Measurements – Model Scattering Matrix**



 $[S_P] = [P_2][S][P_2]^T$ 

- $S \rightarrow$  Scattering Matrix
- Scattering Matrix of Corner Reflector
  S =

 $S = \begin{bmatrix} 1 & 0 \\ 0 & 1 \end{bmatrix}$ 

 $P_{2} \Rightarrow \text{Propagation Matrix} \qquad P_{2} = \begin{bmatrix} \exp(i\kappa_{H}r) & 0 \\ 0 & \exp(i\kappa_{V}r) \end{bmatrix} \qquad \kappa = \frac{2\pi}{\lambda} \left( \cos\theta - \sqrt{\epsilon - \sin^{2}\theta} \right)$  $\epsilon_{HH} \text{ and } \epsilon_{VV} \text{ from snow anisotropy model}$ 

## **Temporal Coherence Region**



• A = 0.2  
• 
$$\rho = 0.2 \frac{g}{cm^3}$$
  
• R1 = 1 cm



## **DInSAR** phase







15





## **Temporal Coherence Region**



270°

DLR



 $\bigcirc$ 

16

270°

## **DInSAR** phase







 Increasing difference between VV and HH

- A1 = 0.1
- A2: anisotropy at 2. acquisition

# 





 Similar behavior as for snow depth change

- A1 = 0.1
- A2: anisotropy at 2. acquisition

## **Temporal Coherence region**



## **Summary and Outlook**

#### Summary

- Modeling of coherence regions for snow depth and anisotropy changes
- Behavior of simulated coherence regions similar to real data
- Phase extent higher sensitivity than CPD

#### **Next Steps**

- Not yet possible to separate anisotropy and snow depth change
  → Further investigation of the influence of snow changes on different polarization states
- Establishment of a retrieval based on coherence region parameters





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**Temporal Coherence Region** 

90°

270°

135°

. 

2250

180°



• A = 0.2  
• 
$$\rho = 0.2 \frac{g}{cm^3}$$
  
• r1 = 1 cm

